Multi-segment earthquakes and the tsunami potential of the Aleutian megathrust

Ian Shennan September 10, 2009

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Spruce trees dying, April 1964 Photo: Gill Mull



Girdwood, Alaska May 2006



Girdwood: 10 peat-mud couplets, 7 great earthquakes



Diatoms in environmental reconstruction







predicted

observed





Girdwood reconstructed land / sea-level change



Last 4000yr: 7 great earthquakes



8

- ~7m net subsidence superimposed on seven earthquake cycles in the past 4000 years.
- EDC model for the Girdwood area with coseismic subsidence (1), followed by rapid postseismic uplift in the decades after the earthquake (2). This merges into centuries of slower inter-seismic uplift (3) before a period of pre-seismic subsidence (4).
- Great earthquakes: no fixed recurrence interval between great. The shortest interval is between ~180 and 720 years. The longest interval is 790 – 920 years, which is between the penultimate earthquake and the Mw9.2 Alaska earthquake of
 9 March 1964.





Zone of co-seismic uplift 1964: Alaganik Slough



Summary

- 1. Net Subsidence (4m since 1500BP)
- 2. Coseismic uplift
- 3. Post-seismic & interseismic subsidence
- Carver & Plafker (2009) summarise 8 earthquakes in 5000yr (correlate with the 4000yr record from Girdwood)



Near the eastern limit of coseismic uplift 1964: Cape Suckling & Bering Glacier



Since the Little Ice Age maximum, ~AD 1900 (dashed line), Bering Glacier has undergone multiple surge and retreat cycles, ~1920, ~1938-1940, 1957-1960, 1965-1967 and 1993-1995.

13 Continued retreat since 1995 exposes new sediment sequences each year.

Cape Suckling marsh







2004 1,2,5,7,8 2005 1-5





Beyond the eastern limit of coseismic uplift 1964: Yakataga to Icy Bay, the Forgotten Coast



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Rapid Communication

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QUATERNARY

Late Quaternary sea-level changes and palaeoseismology of the Bering Glacier region, Alaska

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970±40, 960 to 790BP

26

100









Given the structural setting and historical earthquake history of the region we must two scenarios :

1) a rupture scenario like that of the 1899 Yakataga and 1964 earthquakes, where the Yakataga seismic gap and eastern segment of the Aleutian megathrust ruptured independently and

2) simultaneous rupturing of the megathrust from Cook Inlet in the west to the Pamplona – Malaspina thrust front in the east in which the Yakataga seismic gap ruptures in conjunction with the eastern segment of the Aleutian megathrust.

Two lines of evidence support 2) ~900 BP and ~1500 BP great earthquakes

First, greater coseismic deformation at Cape Suckling in ~900 BP than in 1964 suggests that the area of surface deformation extended farther east.

Second, the more widespread evidence for uplift east of Cape Yakataga in both ~900 BP and ~1500 BP than in 1899



Multi-segment earthquakes

Earthquakes ~900 and ~1500 years ago involved rupture of the western/ central segment of the Yakutat microplate along with the 1964 rupture zone

We calculate an additional 23,000 km² to the rupture area and ~15% increase in seismic moment, from Mw 9.2 in 1964 to ~Mw 9.25 for the multi-segment earthquakes.

Modelling of tsunami generation and coastline inundation for singlesegment and multi-segment earthquakes at other subduction zones suggests that an increased area of sea-floor uplift with a multisegment earthquake produces a tsunami with greater wavelength that penetrates farther inland, even though the height of the wave at the coast may be similar



- Seven great earthquake cycles in the past 4000 years.
- Not a fixed recurrence interval between great, >Mw9, earthquakes. Shortest ~180 – 720 years, longest interval is 790 – 920 years,
- Earthquake Deformation Cycle Model for the area of subsidence: coseismic subsidence, rapid postseismic uplift in the decades after the earthquake; centuries of slower inter-seismic uplift before a period of pre-seismic subsidence.

Multi-segment earthquakes ~900 and ~1500 years ago involved rupture areas ~15% greater than in 1964, indicating greater magnitude, >Mw 9.2, and increased tsunamigenic potential. Complex deformation within the Yakutat microplate, influenced by both subduction and transform fault motion, controls recurrence of these largest earthquakes. Through multiple seismic cycles this leads to net uplift in the Yakataga microplate, part of the Saint Elias orogen.